

Fluvial morphology and sediment transport of the Malekhu Khola, Central Nepal Lesser Himalaya

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ABSTRACT

Areas near the rivers and streams have been widely used for settlement, development works and agriculture due to availability of the resources such as water, aggregates and comparatively easier terrain. It is very important to understand prevailing fluvial conditions for sustainable output. The fluvial conditions of the Malekhu Watershed including the river characteristics, sediment transport and sediment dynamics were studied. Ten river transects and the corresponding segments of the Malekhu Khola were surveyed for cross-sections and longitudinal profiles. Samples were collected in each of the transects and were analysed for suspended sediment concentration. The riverbed sampling was made using Wolman's pebble count for establishing grain size distribution. Manning's roughness coefficients were determined to estimate discharge. The study shows that the Malekhu Khola is a sixth order stream. It has been classified into A4-, B4- and C4-type streams. The hypsometric analysis of the Malekhu Watershed shows that it is in mature stage of erosion. Sediment grain size slightly increases downstream but sorting remains extremely poor to moderately poor. The Malekhu Khola shows eroding tendency at 1.2 km, 16.6 km and 20 km from the origin and brings larger particles only during high flow period at the remaining transects. Mean Suspended Sediment Concentrations (SSC) for the Malekhu Khola was 72.14 mg/L, and it increases with increasing discharge and increasing watershed area and is related to velocity of the river, mining activities and local tributaries.

Key words: Fluvial morphology, hydraulic parameters, hypsometric integral, sediment load, central Nepal

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INTRODUCTION

In the Central Nepal Lesser Himalaya particularly in monsoon season, flooding, erosion, landsliding, etc. occurring in the river basin cause loss of life and property. Streams are dynamic systems that balance water flow and sediment transport. It is known that sediments in the river have their origin from different sources. Therefore, the proportion that each source contributes to the mix varies over time and space, as a result of erosion that is ongoing in the basin of contribution (Poletto et al., 2009). Furthermore, erosion and sedimentation exacerbate the instability of the river channel. This is why it is better to have more

information regarding river; sediment transport, discharge, and dynamics.

Topography, rainfall condition, terrain materials, stream characteristics and landuse play a significant role in the nature and amount of erosion. Singh and Sarangi (2009) performed an assessment of erosion status of two watersheds through hypsometric analysis and highlighted the use of this analysis in interpreting the watershed health condition in terms of erosion. The first attempt to the formulation of a rational theory of bed load transport was made by Du Boys (1879) when he postulated that the bed load discharge should be a function of the difference between the bed shear stress and critical shear stress of the sediment particles in the bed. Understanding of river dynamism and its related processes, and river stability condition are very

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important for watershed management, wetland management, flooding mitigation, sedimentation management and sustainable development of infrastructures (Shrestha and Tamrakar, 2012). Tamrakar and Shrestha (2008) suggested that rivers originating from the Lesser Himalaya with unstable slope are even more capable to flush out sediment compared to the rivers of the Siwaliks. This paper aims to give fluvial morphology, and dynamic scenario of sediment transport by the Malekhu Khola, which lies in the Dhading District of the Central Nepal Lesser Himalaya (Fig. 1).

GEOLOGICAL SETTING

Geology of the study area is based on the work of Stöcklin (1980) and the present survey along the Malekhu Khola section. The Malekhu area lies on the rocks of the Lesser Himalaya, comprising the autochthonous Upper Nawakot Group of the Nawakot Complex, and the Bhimphedi Group and the Tistung Formation (of the Phulchauki Group) of the allochthonous Kathmandu Complex. There is a massive granite intrusion (Agra Granite) in the Tistung Formation of the Phulchauki Group (Table 1; Fig. 2).

The Nawakot Complex comprises the Lower and the Upper Nawakot Groups. The Nawakot complex consists almost exclusively of low-grade metasedimentary rocks. The Kathmandu Complex has been divided into the Precambrian Bhimphedi Group (Nadgir et al. 1968-73), consisting of relatively high-grade metasedimentary rocks, and the Phulchauki Group of sedimentary and weakly metamorphosed rocks of Early-Middle Palaeozoic age. The Mahabharat Thrust passes across the Malekhu Khola separating the Nawakot Complex to the north with the Kathmandu Complex to the south. The dominant structural feature is the large, WNW-ESE trending Mahabharat synclinorium, a doubly plunging megafold with steep flanks, a well-developed western closure and a narrow, elongated eastern wing (Stöcklin, 1980).

METHODOLOGY

The fluvial morphology and sediment transport were included in the study. The fluvial morphology was studied in watershed-scale and reach scale levels. The watershed-scale parameters included stream order, planform, and hypsometric integral. Stream order was identified using Strahler's (1957; 1969) method. Planform of the Malekhu Khola was measured using a toposheets (1:25,000 scale).

Watershed-scale Analysis

The hypsometric integral (HI) is a geomorphological parameter classified under the geologic stages of watershed development. It assumes importance in estimation of erosion status of watershed and subsequent prioritization for taking up soil and water conservation activities. The hypsometric integral is also an indication of the 'cycle of erosion', i.e., time required for reduction of a landtopological unit to the base level (Strahler 1952; Garg 1983). This entire period or the 'cycle of erosion' can be divided into three stages viz. old ($HI < 0.3$), in which the watershed is fully stabilized; equilibrium or mature stage ($0.3 = HI = 0.6$); and inequilibrium or young stage ($HI > 0.6$), in which the watershed is highly susceptible to erosion (Strahler 1952).

For the hypsometric analysis, the subwatershed map along with the DEM of the Malekhu Khola Watershed

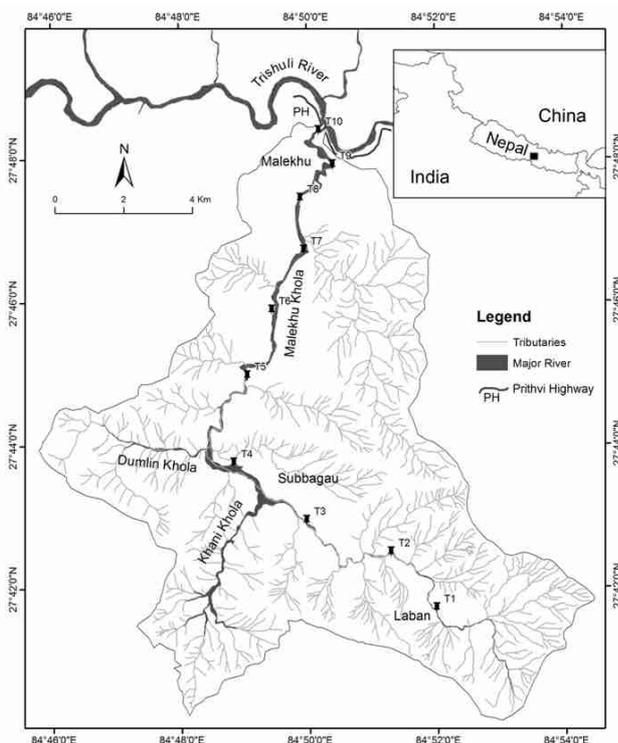


Fig. 1 Location map of the study area

Table Lithostratigraphy of the Geology of the Lesser Himalaya (after Stöcklin, 1980)

Complex	Group	Formation	Main lithology	Thickness (m)	Age	
Kathmandu Complex	Phulchauki	Godawari Limestone	Limestone	300-400	Devonian	
		Chitlang Formation	Slate, Quartzite	1000	Silurian	
		Chandragiri Limestone	Limestone	2000	Cambrian	
		Sopyang Formation	Slate, calc. Phyllite	200	Cambrian (?)	
		Tistung Formation	Metasandstone, Phyllite	300	Early Cambrian or Late Precambrian	
	Bhimphe	Markhu Formation	Marble, schist	1000	Late Precambrian	
		Kulekhani Formation	Quartzite, schist	2000	Precambrian	
		Chisapani Quartzite	White quartzite	400	Precambrian	
		Katitar Formation	Quartzite, schist	2000	Precambrian	
		Bhainsedobhan Marble	Marble	800	Precambrian	
		Raduwa Formation	Garnetiferous schist	1000	Precambrian	
.....Mahabharat Thrust.....						
Nawakot Complex	Upper Nawakot Group	Robang Formation	Phyllite, Quartzite	200-1000	Early Paleozoic	
		Malekhu Limestone	Limestone, Dolomite	800	Early Paleozoic	
		Benighat slate	Slate, argillaceous dolomite	500-3000	Early Paleozoic	
	Unconformity					
	Lower Nawakot Group	Dhading Dolomite	Stromatolitic dolomite	500-1000	Precambrian	
		Nourpul Formation	Phyllite, Metasandstone	800	Precambrian	
		Dandagaun Phyllite	Phyllite	1000	Precambrian	
Fagfog Quartzite		White Quartzite	400	Precambrian		
		Kuncha Formation	Phyllite, Quartzite	3000	Precambrian	

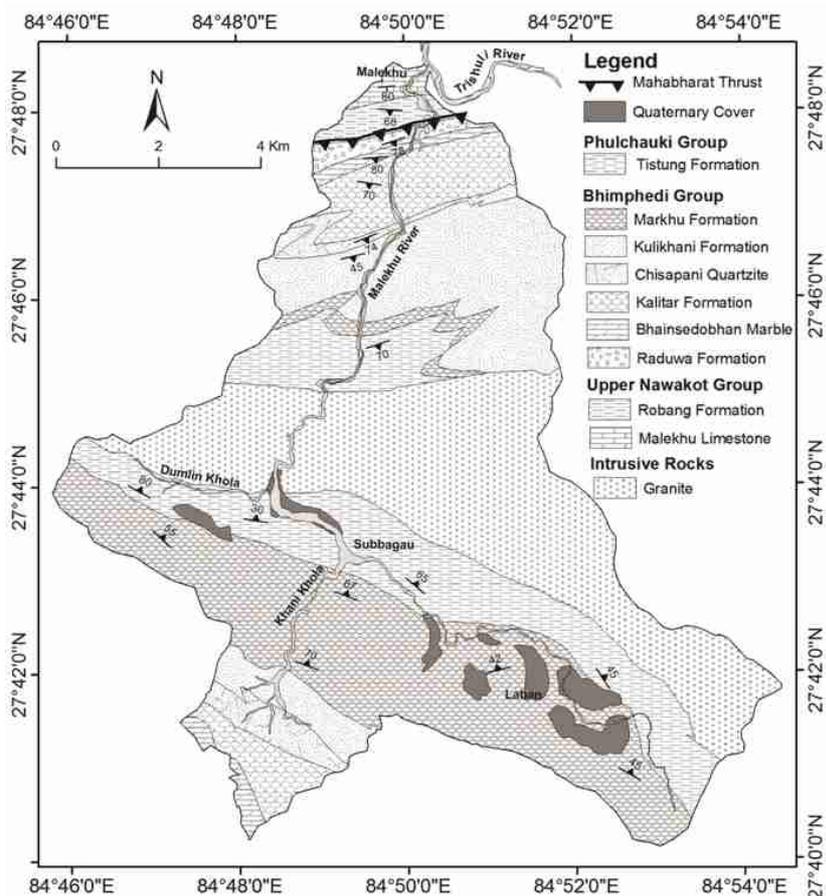


Fig. 2 Geological map of the study area

was prepared in GIS. Then, normalized elevation (h/H) was plotted against the percent area (a/A) for each subwatershed to obtain the hypsometric curve. The hypsometric integral (HI) was obtained by using the following formula proposed by Pike and Wilson (1971):

$$HI = (\text{Elv. mean} - \text{Elv. min.}) / (\text{Elv. max.} - \text{Elv. min.}) \dots (1)$$

Reach-scale Analysis

Cross-section surveys were carried out to generate hydraulic data. Longitudinal profiles were measured for slope. To characterize grain size of each of ten segments in reach-scale and in cross-section scale, Wolman's (1954) pebble counting was adopted. The stream segments were classified into different types designated by Rosgen (1994). Different hydraulic parameters at each transect were plotted against the bankfull area and the subwatershed area contributing upto the transect considered.

The calculation of rate of bed load for the Malekhu Khola was made using the Hassanzadeh (2007) type of common dimensionless formulae on the hydraulics of sediment transport. The existing relations were used to determine the rate of bedload in volume per unit width and unit time (q_s), and rate of bedload in weight per unit time (Q_s). The variables calculated and used in formula are listed in Table 2.

After the calculation of the rate of bedload in volume per unit width and unit time (q_s), the rate of bedload in

weight per unit time (Q_s) was calculated using following formula:

$$Q_s = \gamma_s W_{\text{bkf}} q_s \dots (2)$$

Where, W_{bkf} is the Bankfull width, γ_s is solid particle specific weight.

Most of the suspended sediment concentration (SSC) in a river is due to erosion on the watershed and is not made up of bed material. SSC quantifies concentrations of suspended solid material in surface water. It is usually expressed in milligrams per liter (mg/L). SSC was measured after ASTM D3977-79 (2002) by filtration of a litre of sample, and obtaining the dry weight of all the sediment from a known volume of a water-sediment mixture.

RESULTS

Morphometry

The Malekhu Khola is a 24.83 km long sixth-order stream (Table 3; Fig. 3) and has the watershed area of 101.28 km². The sixth order of the Malekhu Khola is the longest section of length of 11.49 km and also covers largest watershed area of 70.28 km².

Hypsometric Intregal (HI)

Table 2: Dimensionless existing formula to calculate bedload rate (after Hassanzadeh, 2007)

Author	Formula	Explanation
Einstein-Brown (1947)	$q_s / ((\text{ag}d^3))^{0.5} = 23.6 f^3$	Where, a = Immersed sediment specific gravity, $a = (\gamma_s - \gamma) / \gamma$ f = hydrodynamic-immersed gravity force ratio, $f = \{ \gamma R_h S / (\gamma_s - \gamma) d \}$
Bonnefille (1963)	$q_s / ((\text{ag}d^3))^{0.5} = 5.5 f^{1.5} (4.26 f^{0.5} - 1)^{1.25}$	γ = liquid specific weight 1000 kg/m ³ , γ_s = solid particle specific weight 2650 kg/m ³ ,
Hassanzadeh (2007)	$q_s / ((\text{ag}d^3))^{0.5} = 24 f^{2.5}$	R_h = hydraulic radius (m), S = bed slope (m/m), d = particle diameter (d_{50} in m)

Table 3: Stream order, length and subwatershed area of the main stem of the Malekhu Khola

Stream order	1	2	3	4	5	6	Total
Length (km)	0.35	1.16	2.59	6.32	2.92	11.49	27.83
Subwatershed area (km ²)	0.03	0.42	2.35	21.30	6.90	70.28	101.28

There are altogether 91 first order, 20 second order, 27 third order, 3 fourth order and 1 fifth order subwatershed in the whole Malekhu Khola Watershed. The main stem becomes fifth ordered (Strahler's stream order classification basis) when it reaches Talti where it meets the Khani Khola and becomes sixth order river. There are 14 subwatersheds which have HI less than 0.3, 116 subwatersheds which have HI between 0.3 to 0.6 and 12 subwatersheds which have HI greater than 0.6. The first order subwatersheds within the Malekhu Khola Watershed have HI value ranging from 0 to 0.71 (Table 4) while second order subwatersheds have HI ranging from 0.223 to 0.620. The third order subwatersheds have HI ranging from 0.282 to 0.691 and

the fourth order subwatersheds have HI ranging from 0.265 to 0.586. The majority of the subwatersheds are in mature stage (Fig. 4) including the Tudi Khola, Reti Khola, Khahare Khola and Dhobi Khola subwatershed while some of the major tributaries of the Malekhu Khola such as the Khani Khola and Dumlin Khola are in monandnock stage. Some of the subwatersheds in the upstream sections are young in nature and have high erosion characteristics.

Stream Morphology

Table 4: Hypsometric Integral values for the first, second, third and fourth subwatershed of the Malekhu Khola Watershed.

Subwatershed Order	HI value			
	Min.	Max.	Mean	STDEV
1	0	0.71	0.459	0.114
2	0.223	0.62	0.402	0.112
3	0.282	0.691	0.463	0.104
4	0.265	0.586	0.389	0.172

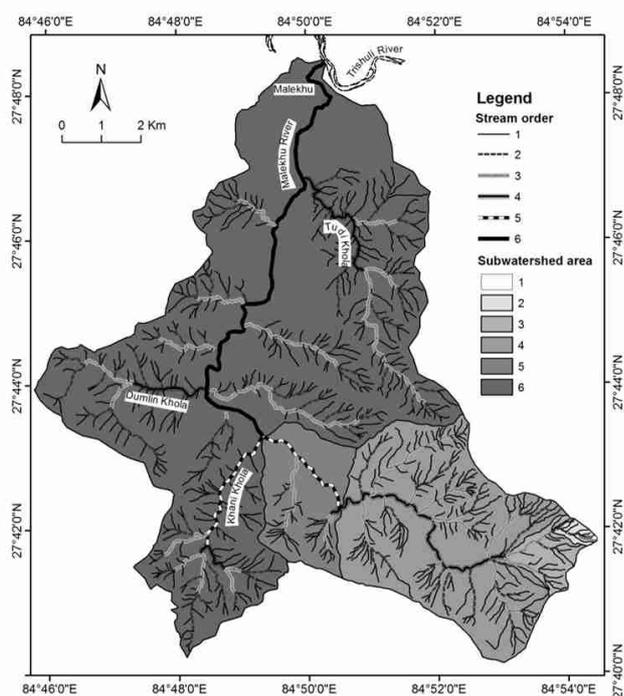


Fig. 3 Stream order and subwatershed map for the main stem and major tributaries of the Malekhu Khola

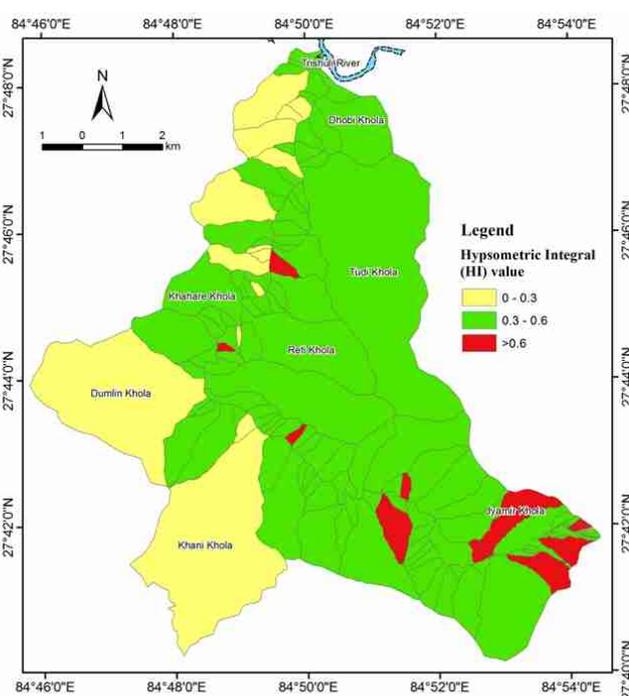


Fig. 4 Hypsometric map of the Malekhu Khola Watershed

Sinuosity (K) ranges from 1.06 to 1.51 (Table 5). T4 and T8 have low sinuosity (K=1.06). This low value of sinuosity indicates that they have very high eroding potential. T5 has the highest value of sinuosity (K=1.51) indicating matured and meandering segment compared to other transects. The meander wavelength ranges from 0.231 to 0.850 km, and the meander belt width ranges from 0.15 to 0.725 km. Both L_m and W_{blt} increase with increasing stream order.

The Malekhu Khola profile is concave upward in shape (Fig. 5). Stream with a steep slope has low sinuosity whereas stream with a gentle slope has high sinuosity. Although slope is partly imposed by geological constraints, streams are nevertheless free to change slope substantially by means of aggradation and degradation.

The Malekhu Khola reaches have been classified as A4-, B4- and C4-type streams (Table 6). A4- and B4-type streams are moderate-gradient gravelly and low sinuosity rivers whereas C4-type streams are low-gradient gravelly meandering river. The Malekhu Khola

gradually changes from A4- to B4-type streams up to 12 km downstream, and then to C4-type stream further downstream, but B4-type streams reappear within the distance of 15 km to 19 km from the origin.

The observed lithofacies on various sedimentary environmental elements are summarized in Table 7. Almost all the transects have dominantly of gravel lithofacies.

Morpho-hydraulic Parameters

The morpho-hydraulic parameters such as bankfull area, bankfull width, bankfull depth, width-depth ratio, floodprone width, entrenchment ratio, wetted perimeter, hydraulic radius, slope, Manning’s n-value, bankfull discharge and bankfull velocity were measured and calculated in field (Table 8). Bankfull cross-sectional area (A_{bkf}) of a stream is the area covered by water during bankfull flow. Bankfull width is the horizontal distance between right and left bankfull point which is measured in field. Entrenchment Ratio (ER) is the ratio of flood prone width to bankfull width. A stream is said

Table 5: Planform parameters of the main stem of the Malekhu Khola

Transects	T1	T2	T3	T4	T5	T6	T7	T8	T9	T10
Sinuosity (K)	1.08	1.15	1.27	1.06	1.51	1.22	1.08	1.07	1.06	1.44
Rc (km)	0.18	0.27	0.16	0.42	0.17	0.31	0.51	0.36	0.07	0.16
L_m (km)	0.68	0.44	0.34	0.76	0.42	0.85	0.84	0.62	0.23	0.51
W_{blt} (km)	0.28	0.15	0.15	0.45	0.53	0.23	0.5	0.33	0.73	0.43
W_{bkf} (m)	11.0	10.0	20.0	30.0	17.5	30.0	18.3	22.3	32.0	23.0
MLR	0.06	0.04	0.02	0.03	0.02	0.03	0.05	0.03	0.01	0.02
MWR	0.03	0.02	0.01	0.02	0.03	0.01	0.03	0.02	0.02	0.02

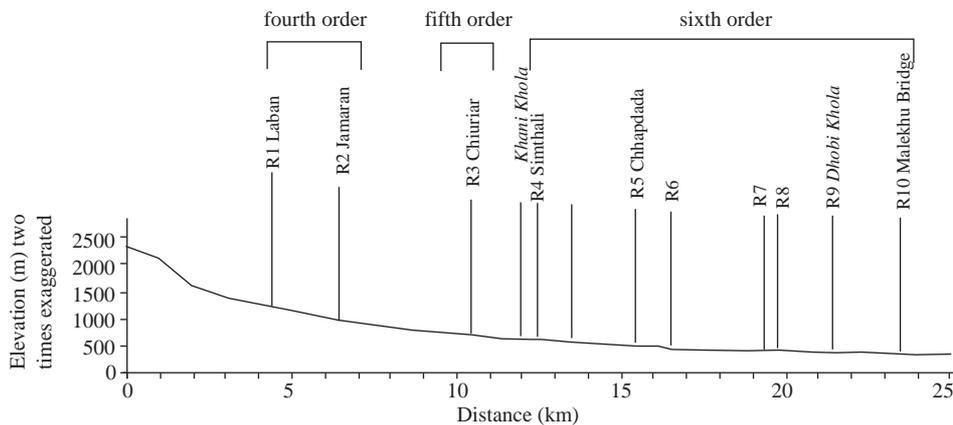


Fig. 5 Longitudinal section of the main stem of the Malekhu Khola

Table 6: Level II classification of the Malekhu Khola

Transect	Location	Stream Order	Geological Formation	Stream Type
1	Laban	4	Markhu Formation	A4
2	Jamaran	4	Markhu Formation	A4
3	Cheuritar	5	Tistung Formation	B4
4	Simthali	6	Tistung Formation	C4
5	Chhpadada	6	Granite Intrusion	B4
6	Kajitol	6	Kulikhani Formation	B4
7	Dwarephat	6	Kulikhani Formation	C4
8	Mawatar	6	Kalitar Formation	C4
9	Dhobi Khola Confluence	6	Robang Formation	C4
10	Malekhu Bridge	6	Malekhu Limestone	C4

Table 7: Major lithofacies observed along Malekhu Khola at ten different transects

Transect	Lithofacies	Remarks
T1 to T2	Channel- cobG, Floodplain-pcobG	Large boulders (b-axis approx 3m) of granite distributed on the main channel, Large angular boulders in floodplain. Presence of large boulders near the channel and reduction of grain size further downstream from the main channel, sparse vegetation on floodplain.
T2 to T3	Channel-pcoG	Thickness of floodplain is about 8 cm in some exposed part and in some area upto 1.5m. Though the floodplain consists of significant amount of boulders and cobbles, it has some area with fine-grained sediments as silt and some clay on its top part. Floodplain is used for agricultural purposes.
T3 to T4	Channel- pcoG and bcoG, Floodplain-cobG, bcopG and pcobG	Effect of the Khani Khola on the lithofacies of channel sediment and floodplain is significant. In some exposed part of the floodplain the thickness is as much as 1 m near Chiuritar. The thickness of the floodplain increases and reaches upto 2 m in exposed part about 150 m downstream from the confluence.
T4 to T5	Channel- bpcog and pcoG, Floodplain- pcobG and cobG, Mid-channel bar- pcoG, Point bar- pcobG	Thickness of floodplain is about 1 to 1.5 m in exposed part. Presence of vegetation on some part of the floodplain while most part was barren with small bushes.
T5 to T6	Channel- copG, Floodplain- pcobG and pcoG, Mid-channel bar- bpcog and pcoG	Thickness of floodplain was about 1m in some exposed part. The vegetation was more on the farthest part of floodplain from main river channel. It was observed that the amount of boulder has been diminished in this region.
T6 to T7	Channel- copG, Floodplain- pcoG and copS, Point bars- pcoG and pS.	Presence of dense vegetation on floodplain. Presence of well sorted sediments on point bars. Thickness of point bar is about 2m.
T7 to T8	Channel- bpcog, Floodplain- bcopG and pcosG, Natural Levee- scoG	Floodplain consists of alternative mud layer and sand layer which are thicker near the channel. The bottom part of the floodplain consists of coarse sediments. Presence of low amount of fine sands on the river bed. Natural levee consists of coarsening upward sequence.
T8 to T9	Channel- scopG, Floodplain- pcoS and bcopS, Point bar- bcopG, Natural Levee- scoG.	Sizes of boulder range from 0.50 to 2m.
T9 to T10	Channel- pcobS, Floodplain- cobG, pcogS, pcoG, copS and scopG.	Presence of chute channel in some upper portions of point bars. The floodplain consists of dispersed coarse-grained sediments such as cobble and boulders probably derived during debris flow.

to be entrenched if the value of entrenchment ratio is less than 1.4 and stream having higher than 1.4 shows affinity to flooding (Rosgen, 1994). Slope of the stream is the ratio of elevation difference to distance of river. Manning’s roughness coefficient (n) depends on the different factors such as channel irregularity, surface roughness, vegetation, seasonal change, channel materials, suspended material, bed material (Cowan, 1956). Due to the lack of the gauge station, velocity (V) and discharge (Q) of the Malekhu Khola were estimated through the Manning's equation and continuity equation (Chow,1959), respectively as:

$$V = (R^{2/3} \cdot S^{1/2}) / n \dots\dots\dots(3)$$

$$Q = V \cdot A_{b_{k_f}} \dots\dots\dots(4)$$

where, R is hydraulic radius (m), S is average channel slope (m/m), n is Manning's roughness coefficient, $A_{b_{k_f}}$ is the cross-sectional area (m²). R is the ratio of $A_{b_{k_f}}$ to wetted perimeter.

The values of the calculated hydraulic parameters (Table 8) are not homogenous for the MRW and their values fluctuate randomly for each transect. Bankfull area is maximum at T9 and minimum at T2. W/D ratio is maximum at T4 and minimum at T1. Flood prone width is maximum at T8 and minimum at T2. Entrenchment ratio is maximum at T8 and minimum at T2. Discharge is maximum at T9 and minimum at T3.

Most of the maximum values are in downstream sections while most of the minimum values are in upstream sections. The geology and the landuse might also be responsible for such fluctuation.

Grain size analysis

The grain size was determined as a textural parameter of the sediments (Table 9). Data obtained from Wolman Pebble counting method was used to calculate grain-size parameters (mean, median, and sorting) by a graphical method (Folk and Ward, 1957).

Median is the diameter corresponding to the 50% mark on the cumulative frequency curve and may be expressed either in phi or in mm. Graphic Mean (M) was calculated from following formula:

$$M = (\phi_{16} + \phi_{50} + \phi_{84}) / 3 \dots\dots(5)$$

Sorting is the measure of degree of scatter. It is the tendency for the grains to all is of one class of grain size. It was computed as:

$$\sigma_I = \{(\phi_{84} - \phi_{16}) / 4\} + \{(\phi_{95} - \phi_5) / 6.6\} \dots\dots(6)$$

It is observed that the sediments along the Malekhu Khola are very poorly sorted on its upstream portion (T1–T5), gradually poorly sorted at T6, and moderately sorted at T7. But at T8 and T9 the sorting of sediment has drastically changed and it is again very poorly

Table 8: Morpho-hydraulic parameters at different transects along the main stem of the Malekhu Khola

Transect	1	2	3	4	5	6	7	8	9	10
$A_{b_{k_f}}$ (m ²)	10.5	4.43	15.32	20.04	11.43	32.67	10.1	27.81	36.33	23.29
$W_{b_{k_f}}$ (m)	11	10	20	30	17.5	30	18.3	22.3	32	23
$D_{b_{k_f}}$ (m)	0.94	0.59	0.74	0.73	0.64	1.2	0.61	1.17	0.87	1.12
W/D	11.7	17.01	27.17	41.1	27.34	25	30	19.06	36.99	20.54
W_{fpa} (m)	22	17	51.5	123	54.5	58	38	129	66	72
ER	2	1.7	2.58	4.1	3.11	1.93	2.08	4.96	2.06	3.13
WP (m)	12.88	11.18	21.47	31.46	18.78	32.4	19.52	24.64	33.73	25.32
R (m)	0.82	0.4	0.71	0.64	0.61	1.01	0.52	1.13	1.07	0.92
S (m/m)	0.03	0.05	0.03	0.04	0.04	0.05	0.03	0.04	0.06	0.05
n-value	0.1	0.08	0.1	0.12	0.14	0.12	0.11	0.09	0.09	0.11
$Q_{b_{k_f}}$ (m ³ /s)	16.59	6.02	4.7	22.85	11.54	62.73	9.76	69.24	97	44.25
$V_{b_{k_f}}$ (m/s)	1.58	1.36	0.31	1.14	1.01	1.92	0.97	2.49	2.67	1.9

$A_{b_{k_f}}$ = Bankfull area, $W_{b_{k_f}}$ = Bankfull width, $D_{b_{k_f}}$ = Bankfull depth, W/D = width/depth ratio, ER = Entrenchment ratio, WP = Wetted perimeter, R = Hydraulic radius, S = Slope, $Q_{b_{k_f}}$ = Bankfull discharge, $V_{b_{k_f}}$ = Bankfull velocity.

Table 9: Grain size distribution at different transects along the main stem of the Malekhu Khola.

	T1		T2		T3		T4		T5		T6		T7		T8		T9		T10	
	Phi	mm	Phi	mm	Phi	mm	Phi	mm	Phi	mm	Phi	mm	Phi	mm	Phi	mm	Phi	mm	Phi	mm
D5	-7	128	-8.6	512	-7.8	256	-11	2048	-7.6	256	-8	256	-6	64	-7	128	-8	256	-8.8	512
D10	-6.8	128	-7	128	-7.2	128	-7.8	256	-7	128	-6.2	64	-5.8	64	-6.6	128	-7	128	-7.8	256
D16	-6.6	128	-6	64	-6.2	64	-6.6	128	-6.4	64	-5.8	64	-5.8	64	-6.2	64	-6.6	128	-7.4	128
D50	-5.2	32	-4.8	32	-5	32	-5	32	-5	32	-5	32	-5.2	32	-4	16	-5.6	64	-5.6	64
D84	-1.4	2	-1	2	-2	4	-2	4	-3	8	-3.8	16	-3.8	16	3.2	0.13	3.2	0.13	-4	16
D95	0.8	0.5	0.6	0.5	1.6	0.25	-0.8	2	3	0.13	-2.6	8	-3	8	3.8	0.06	3.8	0.06	-3	8
M _d	-5.2		-4.8		-5		-5		-5		-5		-5.2		-4		-5.6		-5.6	
M _z	-4.4		-3.9		-4.4		-4.5		-4.8		-4.9		-4.9		-2.3		-3		-5.7	
σ ₁	2.48		2.64		2.47		2.76		2.46		1.32		0.95		3.99		4.24		1.73	

Very well sorted under < 0.35φ; Well sorted 0.35 to 0.50φ; Moderately well sorted 0.50 to 0.71φ; Moderately sorted 0.71 to 1.0φ; Poorly sorted 1.0 to 2.0φ; Very poorly sorted 2.0 to 4.0φ; Extremely poorly sorted Over 4.0φ (After Folk and Ward, 1957)

sorted at T8 and extremely poorly sorted at T9. At T10, channel sediment was found to be poorly sorted. This change in sorting of sediment in its downstream should be the effect of anthropogenic factors. Such factor mainly includes the extraction of certain sized material from river channel as a construction material.

The plot between D₅ with distance (Fig. 6) shows that the value of D₅ slightly decreases with distance. The plot between D₅₀ and distance is observed to change in grain size with distance. It was observed that the maximum grain size occur at T9 and T10. The sorting seems to be poorly related to distance of transport.

Boundary shear stress (τ) and critical shear stress (τ_c) at bankfull stage for d₅₀ were obtained using following expressions respectively (Shields, 1936):

$$\tau = \gamma RS \dots\dots\dots(7)$$

$$\tau_c = \theta_c \cdot g(\rho_c - \rho) \cdot d_{50} \dots\dots(8)$$

where, γ = unit weight of water (9810 N/m²), R = hydraulic radius (m), S = slope of stream (m/), θ_c = Shields parameter for critical dimensionless shear stress which was taken as 0.045 from the Shields Curve for coarse-bedded stream, g = gravitational constant (9.81 m/s²), d₅₀ = median diameter of particle (m), ρ_c = density of sediment (2650 kg/m³), ρ = density of water (1000 kg/m³).

At all transects, boundary shear stress is greater than critical boundary shear stress (Table 9), so the stream is competent enough to carry the bed material distributed on the streambed.

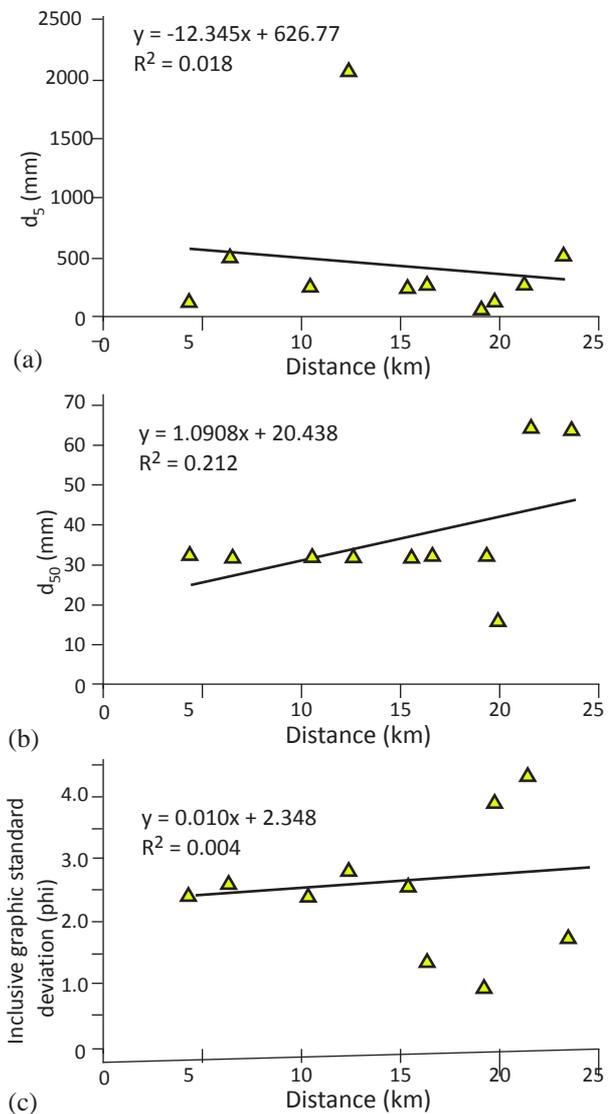


Fig. 6 Grain size parameters against downstream distance. (a) d₅ vs. distance, (b) d₅₀ vs. distance, and (c) Inclusive graphic standard deviation vs. distance

Table 10: Stream competency parameters for the main stem of the Malekhu Khola

Transect	1	2	3	4	5	6	7	8	9	10
d_{50} (m)	0.06	0.06	0.02	0.03	0.03	0.03	0.03	0.03	0.03	0.03
d_{10} (m)	0.128	0.128	0.128	0.256	0.128	0.064	0.064	0.128	0.128	0.256
τ (KN/m ²)	0.241	0.196	0.209	0.251	0.239	0.495	0.153	0.443	0.629	0.451
τ_{c50} (KN/m ²)	0.437	0.437	0.146	0.219	0.219	0.219	0.219	0.219	0.219	0.219
τ_{c10} (KN/m ²)	0.932	0.932	0.932	1.865	0.932	0.466	0.466	0.932	0.932	1.865
Ω (KNm/s/m ²)	0.42	0.60	0.61	0.23	0.26	0.98	0.13	1.05	1.53	0.82

Stream power (Ω) is a measure of the energy transfer (Bagnold, 1966). It is computed by following formula developed by Brookes (1990):

$$\Omega = \tau V \dots\dots\dots (9)$$

where, τ = boundary shear stress (KN/m²), V = bankfull velocity (m/s)

The stream power is maximum at T9 and minimum at T7 (Table 10) along the Malekhu Khola. Since, there is great difference between the boundary shear stress and velocity, the stream power is the function of both boundary shear stress and velocity. The highest value of tractive shear stress at T9 clearly indicates that there is the highest energy or competency to transport bedload. Bedload transport rate obtained from various equations indicates that it is highest at T9. Suspension sediment discharge is also highest at T9. It suggests that the value of tractive shear stress plays an important role in sediment transport. In Malekhu Khola, stream power fluctuates from 0.13 KNm/s/m² to 1.53 KNm/s/m². Stream power is highest at T9 and lowest at T7. The maximum value at T9 is due to the increased velocity (V) which is 2.6 m/s and boundary shear stress (0.629 KN/m²) the lowest value at T7 is due to the low water velocity (0.97 m/s) and boundary shear stress (0.153 KN/m²). Stream power shows huge fluctuation along the stream of transporting capacity influenced probably by change in slope and stream pattern.

Based on plots of hydraulic parameters, the relations of bankfull discharge Q_{bkf} with bankfull area, bankfull width, W_{bkf} and bankfull depth, D_{bkf} are established (Figs. 7a, 7b, 7c). These hydraulic geometry relations indicate positive correlation. Among the relationships,

one of the bankfull area versus bankfull discharge has better relation among the plots.

In the regional curves (Figs. 7d, 7e, 7f), bankfull discharge, bankfull area and bankfull width have been correlated with the drainage area. These relationships show better relations in exponential function and degree of correlation is almost same. With increasing drainage area, the hydraulic parameters tend to increase in exponential function. This indicates that increase stream bankfull cross-sectional area tends to increase bankfull discharge and bankfull width of the Malekhu Khola.

Bedload Transport

Sediment transport characteristics when understood benefits many applications including prediction of effects of landuse and flow regime change and channel restoration efforts (Wilcock, 2001). Sediment transport process is a set of phenomena that mediates between the flowing water and the channel boundary. The sediment load of a river is transported as bedload and/or suspended load and/or wash load. Many researches on sediment transport had been made and extension can be found in Vanoni (1984), Yang (1972; 1973) and Yallin (1963; 1972). The calculation of rate of bed load for the Malekhu Khola was done using the Hassanzadeh (2007) type of common dimensionless formulas on the hydraulics of sediment transport. The existing relations were used to determine the rate of bedload in volume per unit time and unit width (q_s) and rate of bedload in weight per unit time (Q_s). The rate of bed load (Q_s) has though increased (Table 11) from upstream portion of Malekhu Khola towards the downstream portion the trend is fluctuating.

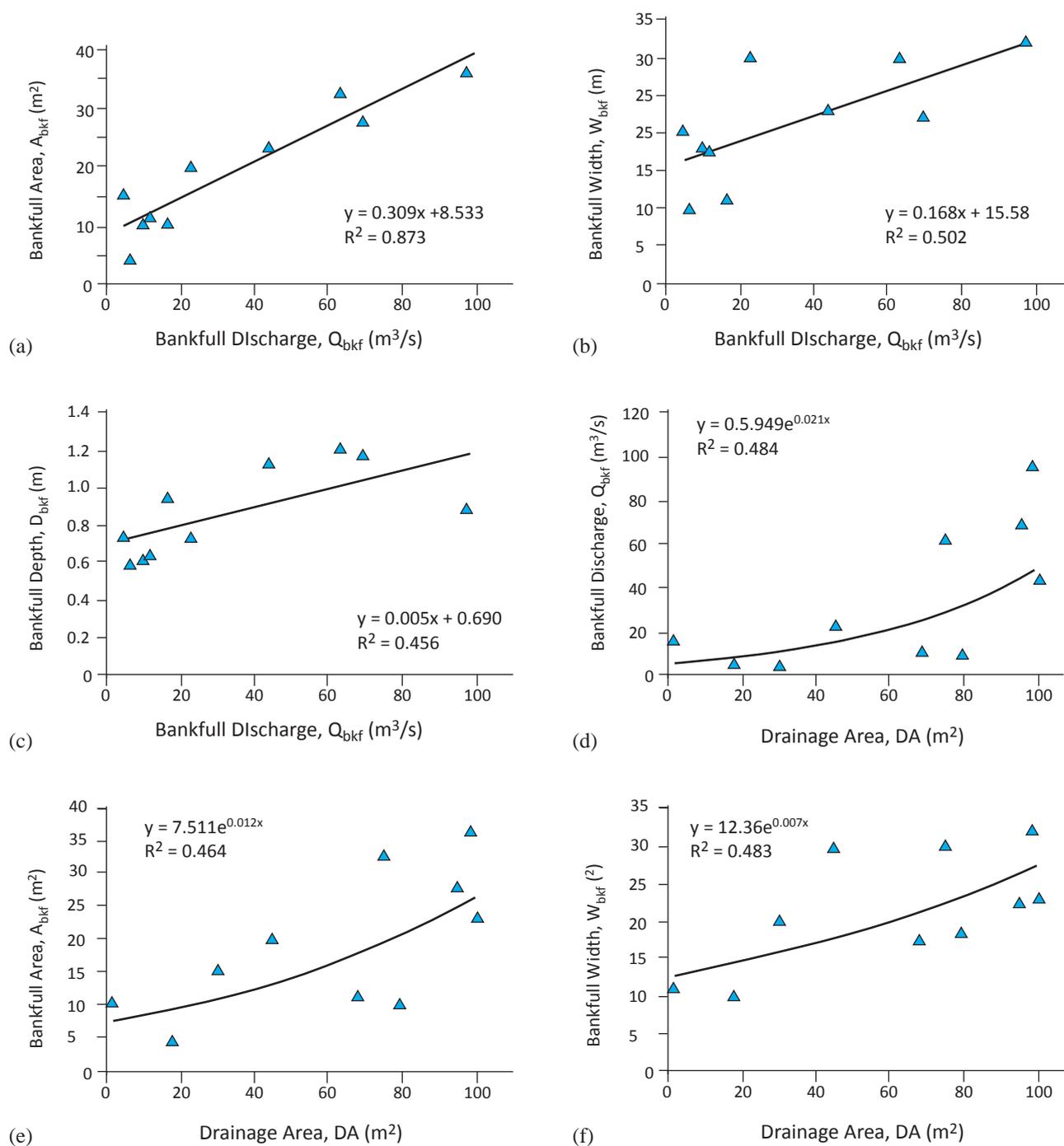


Fig. 7. Hydraulic geometry relations for the Malekhu Khola: (a) Area hydraulic geometry relation, and (b) Width hydraulic geometry relation, (c) Depth hydraulic geometry relation, (d) Discharge regional curve, (e) Area regional curve, and (f) Width regional curve

When the sediment load, i.e., sediment transport capacity in Tonnes/day is plotted against bankfull width (Fig. 8a), all the bedload equations exhibit exponential function of increase of sediment transport capacity of the Malekhu River with increased bankfull width. The

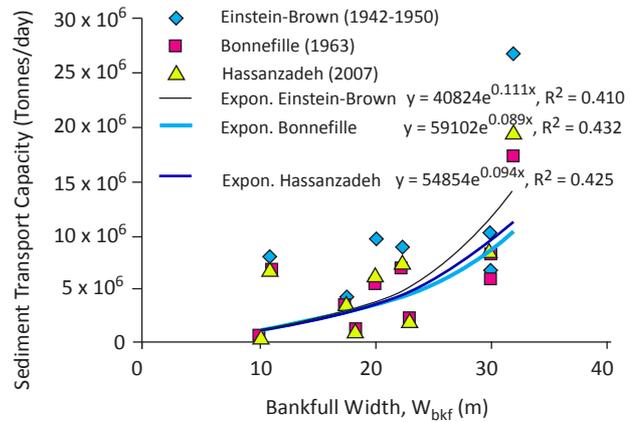
bankfull discharge versus sediment transport capacity curve shows linear function of correlation (Fig. 8b), in which, as discharge increases the sediment transport capacity also tends to increase for the Malekhu Khola.

Table 11: Calculated rate of bedload in weight per unit time Q_s in Tonnes/day using different relations along the Malekhu Khola

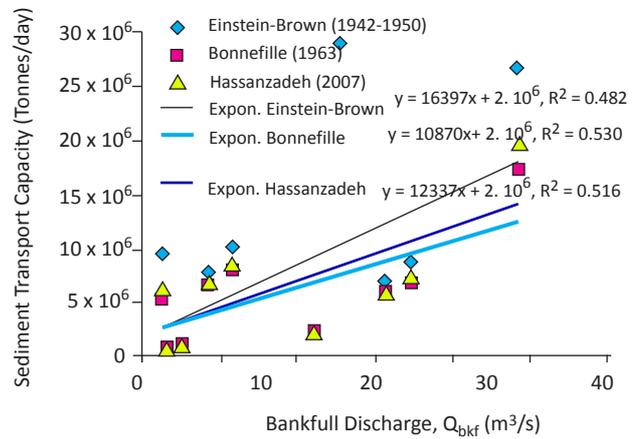
Transects	Bedload transport (Tonnes/day)		
	Einstein and Brown (1942-1950)	Bonnefille (1963)	Hassanzadeh (2007)
T1	7857×10^3	6805×10^3	7101×10^3
T2	290×10^3	495×10^3	447×10^3
T3	9598×10^3	5503×10^3	6554×10^3
T4	10233×10^3	8249×10^3	8796×10^3
T5	3861×10^3	3453×10^3	3569×10^3
T6	6816×10^3	6052×10^3	6270×10^3
T7	938×10^3	1166×10^3	1106×10^3
T8	8903×10^3	6910×10^3	7456×10^3
T9	26657×10^3	17274×10^3	19747×10^3
T10	2129×10^3	2324×10^3	2275×10^3

Suspended Sediment Load

Sediment discharge or sediment load in the channel was calculated by multiplying the main concentration of water at each cross-section by the average discharge. The results of suspended sediment concentration at ten different locations along the Malekhu Khola (Table 12) show that the SSC increases from upstream to downstream. SSC at T9 is exceptionally high (193.78 mg/L) compared to the least value at T8 (20.47 mg/L). The major tributaries of Malekhu Khola has contributed significant amount of suspended sediment to the river system. The Aap Khola is the major tributary contributing maximum suspended sediment which is 997.23 mg/L. The suspended sediment measurement given here is of one time measurement of a particular



(a)



(b)

Fig. 8 (a) Sediment transport capacity vs channel width, and (b) Bankfull discharge versus sediment transport capacity of the Malekhu Khola

Table 12: Suspended sediment discharge of the Malekhu Khola

Transect	SSC (mg/L)	SSC (Kg/m ³)	Suspended Sediment	
			Discharge (kg/s)	Discharge (Tonnes/day)
T1	38.67	0.04	0.64	55
T2	78.33	0.08	0.47	41
T3	40.27	0.04	0.19	16
T4	106.33	0.11	2.43	210
T5	40.42	0.04	0.47	41
T6	84.84	0.08	5.32	460
T7	52.00	0.05	0.51	44
T8	20.47	0.02	1.42	123
T9	193.78	0.19	18.8	1624
T10	66.32	0.07	2.93	253

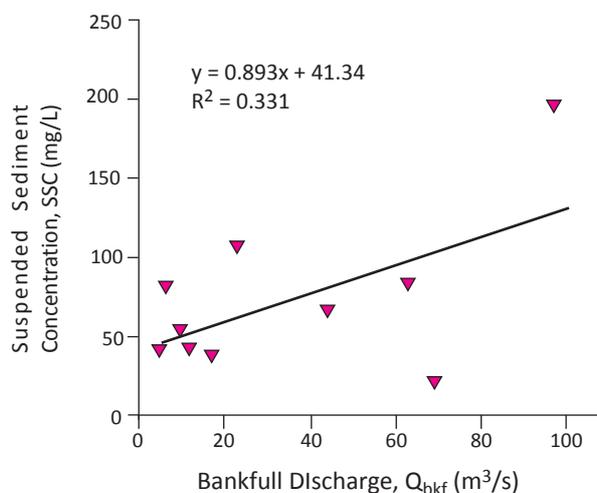


Fig. 9 Sediment rating curve of the Malekhu Khola

season (post monsoon). In this season the sediment rating curve, i.e., the discharge versus suspended sediment concentration (Fig. 9), shows positive but weak correlation due to anomalous result in transect T8.

CONCLUSIONS

The Malekhu Khola is a 6th order stream which is contributed by three major tributaries as 5th order Khani Khola, 4th order Dumling Khola and another 4th order Aap Khola. The majority of the subwatersheds are in mature stage while the big subwatersheds such as those of the Khani Khola and the Dumlin Khola are in relatively old stage. The Malekhu Khola is classified as A4-, B4- and C4- type streams, and is a gravel bed stream. It shows eroding tendency at 1.2 km, 16.6 km and 20 km from the origin, and has enough competency to move larger particles (d_{50}) mostly from the Transect T5 towards the downstream segments. The stream is not competent to carry d_{10} -sized particles at bankfull condition, and therefore these oversized particles are transported only during debris flow.

Sorting of the channel sediments of the Malekhu Khola is extremely poor to moderately poor. The median grain size slightly increases downstream. Sorting and size of channel sediment has been affected by the human influences, i.e., extraction of riverbed materials for construction aggregates.

The hydraulic geometry relations show moderate degree of correlations. The better and positive linear

correlation is found in bankfull cross-sectional area versus bankfull discharge plot. The regional curves showing relationships of hydraulic parameters with the drainage basin area show moderate degree of positive correlation. The hydraulic geometry relations and regional curves are found to be very much influenced and the resulting degree of correlations is low to moderate because of river mining activities.

Bedload transport load calculated ranges between 290×10^3 and 26657×10^3 Tonnes/day. Bedload transport capacity versus bankfull width and bankfull discharge plots also show moderate degree of correlations. Bedload transport rate depends on distance, water discharge, bankfull depth, bankfull width, bankfull velocity and gradient.

Suspended sediment discharge ranges from 55 to 1624 Tonnes/day. The suspended sediment discharge is about 0.02% of the bedload transport, and thus the Malekhu Khola is a bedload river. The sediment rating curve shows poor correlation, and though suspended sediment concentration tends to increase with increased discharge it is much influenced by river mining activities in the downstream segments of the Malekhu Khola.

It is recommended to establish a gauge station to better understand and monitor the sediment transport and discharge of the Malekhu Khola.

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